



## 古盘虫类有孔虫起源及其壳壁发育证据\*

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**摘要** 古盘虫类有孔虫是维宪早期到谢尔普霍夫早期最有用的生物地层标志微体化石之一, 广泛分布于正常浅海底栖环境。古盘虫超科里的平古盘虫亚科从早维宪中期开始出现, 具有最原始的壳壁结构。它将古盘虫类与其祖先假砂盘虫类联系起来。平古盘虫亚科壳壁主要是黑色微粒状。透明放射层局限在脐部或延伸覆盖到内部壳圈的壳缘。平古盘虫从何而来? 即古盘虫类的祖先问题, 一直存在争议。本文通过研究湖南邵阳马栏边剖面石磴子组连续演化的*Lapparentidiscus*和*Viseidiscus*, 证实了*Lapparentidiscus*是古盘虫类的祖先, 早期演化序列为*Lapparentidiscus talasicus*–*Viseidiscus eospirillinoides*–*Viseidiscus monstratus*, 古盘虫类的透明放射层是以替代而非填充的形式出现的, 透明放射层逐渐替代黑色微粒层, 并从脐部逐步扩展到壳缘。

**关键词** 有孔虫 古盘虫类 维宪阶 华南 演化序列

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## The origin of primitive archaediscacean foraminifers and the development of their wall structure

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**Abstract** Archaediscacean foraminifers are perhaps one of the most useful biostratigraphic marker microfossils from the early Viséan to early Serpukhovian. They are widespread in normal marine, benthic, shallow-shelf carbonate environments. The subfamily Planoarchaediscinae of the superfamily Archaediscacea appeared in the middle early Viséan. The wall structure of the planoarchaediscins is the most primitive in this superfamily, which provides an evolutionary link between the Archaediscidae and the ancestral pseudoammodiscids. The wall of the archaediscaceans is predominantly dark and microgranular. The hyaline-radial layer is either limited to the umbilical area, or it forms a thin covering over the sides, or sometimes it covers the periphery of the interior volutions. Regarding the ancestor of the archaediscids, this topic is still open to debate. This study focuses on the specimens of *Lapparentidiscus* and *Viseidiscus* with continuous stratigraphic appearances from the Shidengzi Formation at the Malanbian section in South China. The result suggests that *Lapparentidiscus* is a unique pseudoammodiscid that gives rise to the archaediscids. The evolution trend of primitive archaediscaceans is from dark and microgranular to hyaline and granular.

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cean foraminifers is: *Lapparentidiscus talasicus*–*Viseidiscus eospirillinooides*–*Viseidiscus monstratus*. The hyaline-radial layer of the archaetidiscids appears as a replacement instead of a filling. It replaces the dark, microgranular layer gradually, and extends from the umbilical area to the sides.

## SYSTEMATIC PALEONTOLOGY

**Order FORAMINIFERA** Eichwald, 1830

**Suborder FUSULINIDA** Wedekind, 1937

**Family Pseudoammodiscidae** Conil and Lys in Conil and Pirlet, 1970

**Genus *Lapparentidiscus*** Vachard, 1980

***Lapparentidiscus talasicus* (Mikhno in Mikhno and Balakin, 1975)**

(Fig. 2-1–2-7)

**Diagnosis** The test is small, narrowly discoidal, and shallowly umbilicate. The undivided, planispirally coiled tubular chamber comprises 4 to 5.5 whorls. Coiling is initially involute and evolute in the last one or two whorls. The chamber, circular in axial section, expands very quickly in the last few whorls. The wall is dark and microgranular. The aperture is simple and terminal. The diameter of the test is 250–400  $\mu\text{m}$ ; the width of the test is 60–100  $\mu\text{m}$ ; and the width to diameter ratio is 0.20–0.25. The diameter of the proloculus is 15–30  $\mu\text{m}$ . The height of the tubular chamber is 40–80  $\mu\text{m}$  in the last whorl. The thickness of the wall of the last whorl is 12–16  $\mu\text{m}$ .

**Locality and horizon** Malanbian section, Shaoyang County, Hunan Province; 334.3–347 m of the Shidengzi Formation.

**Superfamily Archaetidiscacea** Cushman, 1928 *nom. trans.* Piller, 1978

**Family Archaetidiscidae** Cushman, 1928 *nom. trans.* Chernysheva, 1948

**Subfamily Planoarchaetidiscinae** Mamet, 1975

**Genus *Viseidiscus*** Mamet, 1975

***Viseidiscus eospirillinooides* (Brazhnikova, 1967)**

(Fig. 2-8–2-14)

**Diagnosis** The test is small, narrowly discoidal, and shallowly umbilicate. The undivided, planispirally coiled tubular chamber comprises 4 to 5 whorls. Coiling is

initially involute and subsequently evolute. The chamber expands relatively quickly in the last few whorls. The wall is predominantly dark and microgranular. The weak, unstable, discontinuous hyaline-radial layer is limited to the umbilical area. The aperture is simple and terminal. The diameter of the test is 240–350  $\mu\text{m}$ ; the width of the test is 60–95  $\mu\text{m}$ ; and the width to diameter ratio is 0.22–0.30. The diameter of the proloculus is 14–20  $\mu\text{m}$ . The height of the tubular chamber is 45–70  $\mu\text{m}$  in the last whorl. The thickness of the wall of the last whorl is 10–15  $\mu\text{m}$ .

**Locality and horizon** Malanbian section, Shaoyang County, Hunan Province; 339–344 m of the Shidengzi Formation.

***Viseidiscus monstratus* (Grozdilova and Lebedeva, 1954)**

(Fig. 2-15–2-20)

**Diagnosis** The test is small, narrowly discoidal, and shallowly umbilicate. The undivided, planispirally coiled tubular chamber comprises 4 to 5 whorls. Coiling is initially involute and evolute in the last one or two whorls. The lumina is open. The chamber floor is convex. The chamber expands rapidly in the last few whorls. The wall is composed of a continuous dark microgranular layer and a thin hyaline-radial layer deposited on umbilical sides and, in some cases, also over the periphery of initial volutions. The hyaline-radial layer is absent in evolute whorls. The aperture is simple and terminal. The diameter of the test is 200–350  $\mu\text{m}$ ; the width of the test is 60–90  $\mu\text{m}$  and the width to diameter ratio is 0.20–0.27. The diameter of the proloculus is 15–25  $\mu\text{m}$ . The height of the tubular chamber is 45–70  $\mu\text{m}$  in the last whorl. The thickness of the wall of the last whorl is 10–15  $\mu\text{m}$ .

**Locality and horizon** Malanbian section, Shaoyang County, Hunan Province; 340.2–347 m of the Shidengzi Formation.

**Key words** foraminifera, archaetidiscids, Visean, South China, evolution pattern

## 1 前 言

古盘虫类有孔虫最早出现在石炭纪维宪早期,在正常浅海底栖环境下广泛分布,直到宾夕法尼亚亚纪中期消失。它们是维宪早期到谢尔普霍夫早期最典型的生物地层标志微体化石之一。古盘虫最基本的形态特征是发育初房和没有隔壁的、绕卷的管状第二房室。壳壁结构、壳体形状、绕卷方式、缝合线形态和次生沉积的发育状况形

成了丰富的形态特征,也是分类的重要依据。关于古盘虫类的分类体系存在很多的争论(Cushman, 1928; Chernysheva, 1948; Dain and Grozdilova, 1953; Miklukho-Maklay, 1957; Mamet, 1975; Pirlet and Conil, 1977; Loeblich and Tappan, 1984; Brenckle *et al.*, 1987; Vachard, 1988)。Brenckle等(1987)的分类方案(表1)反映了两个主要的演化趋势:壳壁结构中微粒层的消减和管状房室的充填(图1)。

表 1 古盘虫类有孔虫的分类方案(修改自 Brenckle *et al.*, 1987)  
Table 1 Classification of the archaediscaean foraminifers (modified from Brenckle *et al.*, 1987)

古盘虫超科 Archaediscaea (部分)			
古盘虫科 Archaediscaidae			星古盘虫科 Asteroarchaediscaidae
平古盘虫亚科 Planoarchaediscainae	哈萨克斯坦盘虫亚科 Kasachstanodiscinae	古盘虫亚科 Archaediscainae	
<i>Viseidiscus</i>	<i>Glomodiscus</i>	<i>Archaediscus</i>	<i>Asteroarchaediscus</i>
<i>Planoarchaediscaus</i>	<i>Uralodiscus</i>	<i>Planospirodiscus</i>	<i>Neoarchaediscaus</i>
	<i>Paraarchaediscaus</i>	<i>Betpakodiscus</i>	<i>Permodiscus</i>
	<i>Kasachstanodiscus</i>	<i>Browndiscus</i>	<i>Brenckleina</i>
		<i>Eosigmoilina</i>	

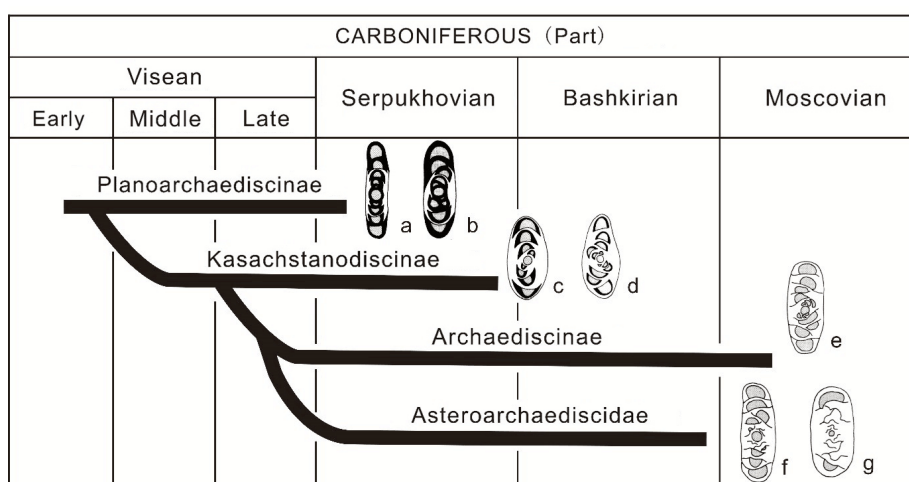


图 1 古盘虫超科的演化(各亚科起止时代来自 Brenckle *et al.*, 1987)

Fig. 1 Phylogeny of the Archaediscaea (ranges of subfamilies are from Brenckle *et al.*, 1987)

素描图为样品中出现的属: a. *Viseidiscus*; b. *Planoarchaediscaus*; c. *Glomodiscus*; d. *Paraarchaediscaus*; e. *Archaediscus*; f. *Neoarchaediscaus*; g. *Asteroarchaediscaus*。

Illustrated genera: a. *Viseidiscus*; b. *Planoarchaediscaus*; c. *Glomodiscus*; d. *Paraarchaediscaus*; e. *Archaediscus*; f. *Neoarchaediscaus*; g. *Asteroarchaediscaus*。

壳壁结构中微粒层的消减体现在古盘虫超科的起源与早期演化中。古盘虫超科里平古盘虫亚科从早维宪中期开始出现, 它的壳壁结构是最原始的。它将古盘虫类与其祖先假砂盘虫类联系起来。平古盘虫亚科壳壁主要是黑色微粒状。透明放射层局限在脐部或延伸覆盖到内部壳圈的壳缘。该亚科有两个属: *Viseidiscus* 和 *Planoarchaediscaus*, 其中 *Viseidiscus* 包括所有壳壁以黑色微粒层为主的平旋的古盘虫类, *Planoarchaediscaus* 则包括其他壳壁以黑色微粒层为主、早期倾斜绕卷、后期平旋的古盘虫类。哈萨克斯坦盘虫亚科在接近维宪中期时开始出现, 黑色微粒层和透明放射层都比较发育。这种壳壁结构反映了平古盘虫亚科到古盘虫亚科转变

的过渡阶段, 即黑色微粒层为主过渡到透明放射层为主。在哈萨克斯坦盘虫亚科内, 也体现了这种壳壁结构的转变, *Glomodiscus* 和 *Uralodiscus* 黑色微粒层非常发育, 而到了 *Kasachstanodiscus* 和 *Paraarchaediscaus* 黑色微粒层明显变弱了。

古盘虫亚科从维宪晚期开始出现, 它的透明放射层是古盘虫科里最发育的。至星古盘虫科, 壳壁的黑色微粒层仅在少数标本中发育。这些标本主要发现在维宪晚期到谢尔普霍夫早期。这一时期正是哈萨克斯坦盘虫亚科逐渐演化成古盘虫亚科的阶段。星古盘虫科从维宪最晚期开始出现, 管状房室部分或几乎全部被有机成因的次生沉积充填。壳壁只有一层透明放射层。

关于古盘虫类的祖先, Conil等(1980)认为 *Brunsia* 和 *Pseudoammodiscus* 是古盘虫类的两个祖先。但是, Hance等(2011)认为 *Lapparentidiscus* 是古盘虫类的唯一祖先。Vachard (1988)指出古盘虫类的透明放射层是以替代而非填充的形式出现的。这些争议一直没有解决。华南古盘虫类有孔虫极为发育。本文试图通过研究华南的早期古盘虫类, 来探讨古盘虫类的祖先和早期的演化序列。

## 2 研究材料与剖面概况

本文的研究材料采自湖南邵阳的马栏边剖面(起点处GPS: 27°22'35" N、111°33'40" E)。该剖面位于湖南省邵阳市新邵县田心乡马栏边水库旁, 在长沙西南方向约170 km, 隶属于湘桂地层小区。湘桂地层小区位于扬子陆块与华夏陆块之间的陆间裂谷盆地; 由东、西两侧NE、NEE向断裂控制, 南与钦州盆地相接, 包括湘南、湘中、桂中和粤西等地(Wang *et al.*, 2013)。密西西比亚纪早期沉积以碳酸盐岩为主, 夹少量砂泥质碎屑岩; 密西西比亚纪晚期沉积三分, 上、下部为灰岩, 中部是含煤碎屑岩系, 部分地区上部是灰岩夹石膏层, 为局限沉积环境。宾夕法尼亚亚纪本区普遍发育碳酸盐岩沉积, 为较深水缓坡、台地沉积环境, 局部地区为局限、浅滩沉积环境(Wang *et al.*, 2013)。

马栏边剖面发育了一套泥盆系到石炭系的地层, 是湘中泥盆-石炭系的经典剖面, 其沉积、生物和同位素地层都有过研究(裘松余、林甲兴, 1982; 湖南省地质矿产局区域地质调查队, 1987; 谭正修, 1990; Hance *et al.*, 2011; Yao *et al.*, 2015)。本文讨论的有孔虫化石来自马栏边剖面石磴子组最上部的一个有孔虫带 *Viseidiscus* 带(盛青怡, 2016), 这是维宪阶的第二个有孔虫带(Wang *et al.*, 2019)。石磴子组下伏陡岭坳组, 上覆测水组, 厚152 m, 在剖面中自195–347 m, 岩性以厚层灰岩为主, 地层出露连续, 无明显的沉积间断, 以有孔虫为代表的浅水相化石丰富, 时代从杜内晚期到维宪早期, 是开展古盘虫类早期演化序列研究的理想层段。

## 3 系统古生物学

拉伯盘虫属 Genus *Lapparentidiscus* Vachard, 1980

塔拉斯拉伯盘虫 *Lapparentidiscus talasicus* (Mikhno in Mikhno and Balakin, 1975)

(图2-1-2-7)

1975 *Ammodiscus?* *talasicus* Mikhno in Mikhno and Balakin, p. 28, 29, pl. 2, figs. 11–13.

1980 *Lapparentidiscus bokanensis* Vachard, p. 285–290, pl. 15, figs. 1, 2.

1984 *Eodiscus hubeiensis* 林甲兴, 141页, 图版6, 图31–33。

1990 *Quasidiscus hubeiensis* 林甲等, 205页, 图版22, 图28, 29。

2011 *Lapparentidiscus hubeiensis* Hance *et al.*, p. 47, pl. 44, figs. 6–8, 10–12.

**特征** 壳小, 窄形圆盘状, 脐部微凹。管状房室无隔壁, 平旋。4–5.5个壳圈, 大部分内旋, 只有最后一、两个壳圈外旋。最后一两圈扩展速度明显加快, 在轴切面呈近圆形。壳壁黑色、微粒状。口孔简单, 位于房室末端。壳体直径为250–400  $\mu\text{m}$ , 宽度为60–100  $\mu\text{m}$ , 轴率为0.20–0.25。初房直径为15–30  $\mu\text{m}$ 。最后一圈房室高度40–80  $\mu\text{m}$ 。最后一圈壳壁厚度为12–16  $\mu\text{m}$ 。

**讨论** Vachard (1980)在建立 *Lapparentidiscus* 属时命名了新种 *L. bokanensis* 作为这个属的模式种。但是, *L. bokanensis* 其实是 *Ammodiscus?* *talasicus* Mikhno in Mikhno and Balakin, 1975的同物异名。*L. bokanensis* 的模式标本来自阿富汗中部的  $V_{2a}$ 带(维宪中下部), *L. talasicus* 的模式标本来自于西天山的  $C_{IV1}$ 层(维宪下部)。*Lapparentidiscus* 在文献里比较少见, 主要因为有些标本被放在 *Pseudoammodiscus* 或 *Mediocris* 这两个属里了。在中国曾被放在 *Eodiscus* Vdovenko, 1970和 *Quasidiscus* Lin *et al.*, 1990这两个属里, 前者建属时没考虑到这个属名已经被三叶虫类占用, 林甲等(1990)建立新属名 *Quasidiscus* 来代替 *Eodiscus*, 但是 *Lapparentidiscus* Vachard, 1980建立时间在前。根据优先律, 应用属名 *Lapparentidiscus* Vachard, 1980代替 *Quasidiscus* Lin *et al.*, 1990。

**产地层位** 湖南邵阳马栏边剖面石磴子组334.3–347 m (图3)。

**时代分布** 维宪期。主要分布在古特提斯地区, 现俄罗斯、乌克兰、法国、伊朗、乌兹别克

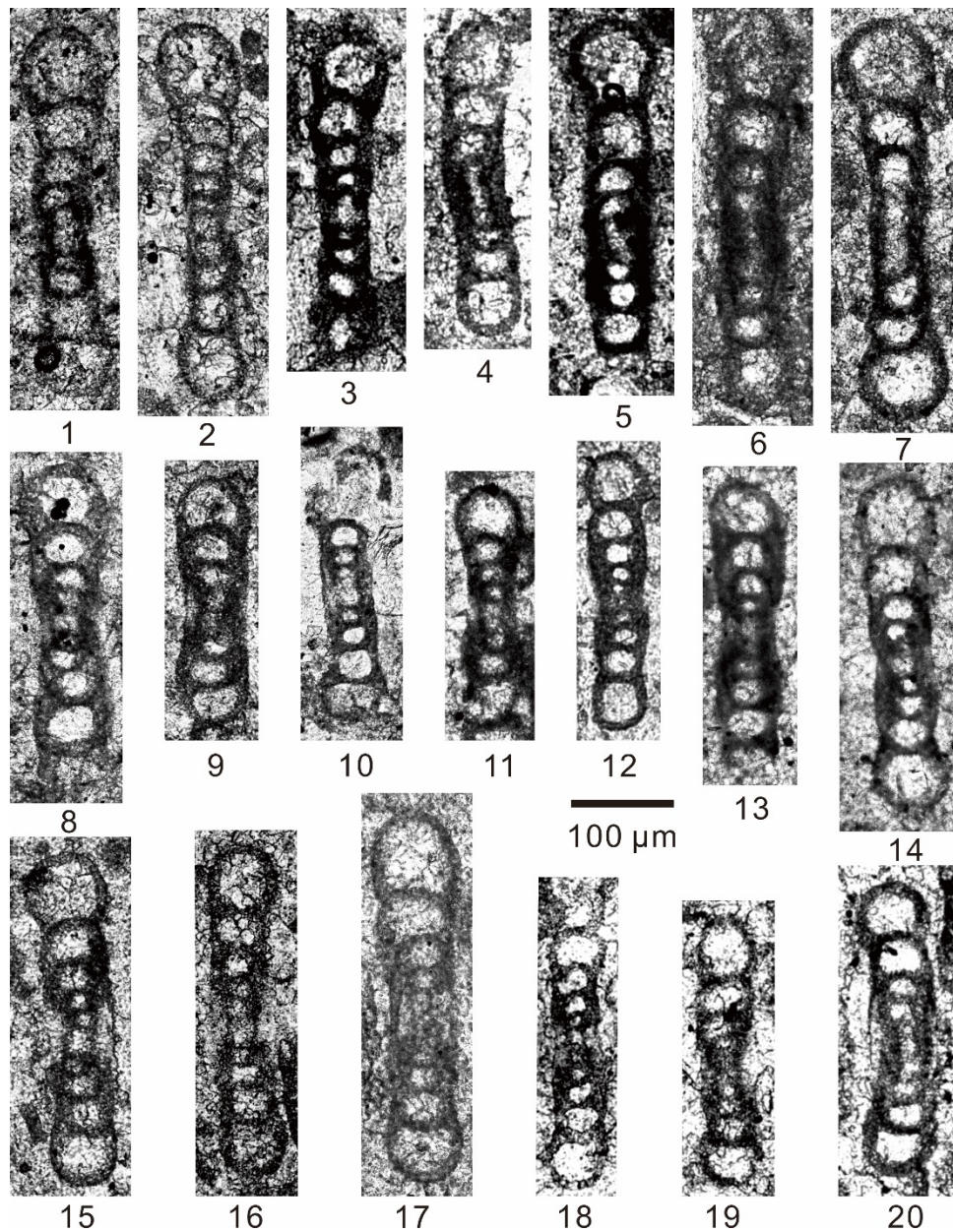


图 2 马栏边剖面的拉伯盘虫和威赛盘虫

Fig. 2 *Lapparentidiscus* and *Viseidiscus* from the Malanbian section

1-7. *Lapparentidiscus talasicus* (Mikhno in Mikhno and Balakin, 1975): 1. 轴切面, 石磴子组 334.3 m (标本号 3340301); 2. 轴切面, 340.3 m (3400301); 3. 轴切面, 341 m (3410001); 4. 近轴切面, 342 m (3420001); 5. 轴切面, 341 m (3410002); 6. 近轴切面, 344 m (3440001); 7. 近轴切面, 340.3 m (3400302); 8-14. *Viseidiscus eospirillinoides* (Brazhnikova, 1967): 8. 轴切面, 339.3 m (3390301); 9. 近轴切面, 339 m (3390001); 10. 轴切面, 340.3 m (3400303); 11. 轴切面, 340.6 m (3400601); 12. 轴切面, 340.2 m (3400201); 13. 轴切面, 344 m (3440002); 14. 轴切面, 340.6 m (3400602); 15-20. *Viseidiscus monstratus* (Grozdilova and Lebedeva, 1954): 15. 轴切面, 340.6 m (3400603); 16. 轴切面, 340.3 m (3400304); 17. 近轴切面, 347 m (3470001); 18. 轴切面, 340.2 m (3400202); 19. 近轴切面, 343 m (3430001); 20. 轴切面, 343 m (3430002).

1-7. *Lapparentidiscus talasicus* (Mikhno in Mikhno and Balakin, 1975): 1. axial section, Shidengzi Formation, 334.3 m (No. 3340301); 2. axial section, 340.3 m (3400301); 3. axial section, 341 m (3410001); 4. near-axial section, 342 m (3420001); 5. axial section, 341 m (3410002); 6. near-axial section, 344 m (3440001); 7. near-axial section, 340.3 m (3400302); 8-14. *Viseidiscus eospirillinoides* (Brazhnikova, 1967): 8. axial section, 339.3 m (3390301); 9. near-axial section, 339 m (3390001); 10. axial section, 340.3 m (3400303); 11. axial section, 340.6 m (3400601); 12. axial section, 340.2 m (3400201); 13. axial section, 344 m (3440002); 14. axial section, 340.6 m (3400602); 15-20. *Viseidiscus monstratus* (Grozdilova and Lebedeva, 1954): 15. axial section, 340.6 m (3400603); 16. axial section, 340.3 m (3400304); 17. near-axial section, 347 m (3470001); 18. axial section, 340.2 m (3400202); 19. near-axial section, 343 m (3430001); 20. axial section, 343 m (3430002).

斯坦、阿富汗、摩洛哥、土耳其等国家境内, 以及中国南部和塔里木等地区。

### 威赛盘虫属 Genus *Viseidiscus* Mamet, 1975

#### 始绕旋威赛盘虫 *Viseidiscus eospirillinoides* (Brazhnikova, 1967)

(图2-8-2-14)

1967 *Planoarchaediscus eospirillinoides* forma *typica* Brazhnikova in Brazhnikova *et al.*, p. 164, 165, pl. 10, fig. 10, pl. 51, figs. 7-9, 12, 13.

1967 *Planoarchaediscus eospirillinoides* forma *involuta* Brazhnikova in Brazhnikova *et al.*, p. 165, 166, pl. 51, figs. 10, 11, 14.

1983 *Planoarchaediscus eospirillinoides* 王克良, 图版3, 图10, 11。

2016 *Lapparentidiscus bokanensis* Zandkarimi *et al.*, fig. 6, no. 25.

**特征** 壳体小。圆盘状, 脐部微凹。管状房室无隔壁, 平旋。4-5个壳圈。大部分壳圈内旋, 只有最后一、两个壳圈外旋。壳圈的扩张速度比较稳定, 除了最后一两个壳圈扩张速度较快。壳壁黑色、微粒状, 只在脐部出现微弱的、不稳定、不连续的透明放射层。口孔简单, 房室末端。壳体直径为 240-350  $\mu\text{m}$ , 宽度为 60-95  $\mu\text{m}$ , 轴率为 0.22-0.30。初房直径为 14-20  $\mu\text{m}$ 。最后一圈房室高度为 45-70  $\mu\text{m}$ 。最后一圈壳壁厚度为 10-15  $\mu\text{m}$ 。

**讨论** 模式标本来自于第聂伯罗-顿涅茨凹地 (Dnieper-Donets Depression) 维宪阶 Yasnaya Polyana 亚阶的 Tula 层。Brazhnikova (1967) 在命名新种的时候, 将 *P. eospirillinoides* 分成两个变型: *typicus* 和 *involutus*。Brenckle 和 Marchant (1987) 认为后一个变型是前一个的幼年体, 没必要分出来。本文认同这种观点, 在观察化石薄片的过程中, 也发现了幼年体 (3-4 个壳圈) 与成年体同时存在的现象。Brazhnikova (1967) 在描述 *P. eospirillinoides* 时提到, 产自 Tula 层的 *P. eospirillinoides* 与产自之上较新的 Oka 亚阶的 *P. monstratus* 形态相近, 两者最大的区别就是壳壁结构, *P. eospirillinoides* 黑色微粒壳壁中有不连续的透明放射颗粒。

**产地层位** 湖南邵阳马栏边剖面石磴子组 339-344 m (图3)。

**时代分布** 维宪期。乌克兰、伊朗、捷克和中国南部地区。

#### 蒙斯特拉威赛盘虫 *Viseidiscus monstratus* (Grozdilova and Lebedeva, 1954)

(图2-15-2-20)

1954 *Archaediscus monstratus* Grozdilova and Lebedeva, p. 61, pl. 7,

figs. 17, 18.

1966 *Planoarchaediscus monstratus* Bogush and Yuferev, p. 159, pl. 11, figs. 4, 5.

1970 *Planoarchaediscus monstratus* Bogush and Yuferev, pl. 8, figs. 14, 15.

1976 *Ammarchaediscus* (*Ammarchaediscus*) sp. Conil, pl. 3, fig. 26.

1977 *Ammarchaediscus* (*Ammarchaediscus*) *monstratus* Conil and Pirlet in Pirlet and Conil, pl. 3, figs. 53-56.

1978 *Viseidiscus kumyrlensis* Marfenkova, p. 98, pl. 8, fig. 21.

1980 *Ammarchaediscus* (*Ammarchaediscus*) *monstratus* Conil *et al.*, pl. 7, figs. 17, 18 (= Pirlet and Conil, 1977, pl. 3, figs. 55, 56).

1987 *Viseidiscus monstratus* Brenckle and Marchant, p. 81, pl. 2, figs. 1-7.

2003 *Viseidiscus monstratus* Brenckle and Milkina, pl. 2, fig. 9.

2009 *Viseidiscus monstratus* Brenckle *et al.*, pl. 8, figs. 11, 18-20.

2011 *Ammarchaediscus involutus* Hance *et al.*, p. 49, 50, pl. 51, fig. 4.

2015 *Viseidiscus monstratus* Shen and Wang, fig. 4, A, C.

2016 *Ammarchaediscus* cf. *eospirillinoides* Zandkarimi *et al.*, fig. 6, no. 29.

**描述** 壳体小, 窄形圆盘状, 脐部宽。管状房室无隔壁, 平旋。4-5个壳圈。大部分壳圈内旋, 只有最后一、两个壳圈外旋。房室无充填。房室底面凸。口孔简单, 位于管状房室末端。壳壁主要由连续的黑色微粒层组成, 透明放射层局限在脐部, 或有时延伸覆盖到内部壳圈的壳缘。壳体后期外旋的壳圈没有透明放射层。壳圈数: 4-5; 直径: 200-350  $\mu\text{m}$ ; 宽度: 60-90  $\mu\text{m}$ ; 轴率: 0.20-0.27; 初房直径: 15-25  $\mu\text{m}$ ; 最后一圈壳壁厚度: 10-15  $\mu\text{m}$ ; 最后一圈房室高度: 45-70  $\mu\text{m}$ 。

**讨论** 模式标本来自于俄罗斯维宪阶的 Oka 亚阶。Grozdilova 和 Lebedeva (1954) 在描述模式种的时候, 说早期的壳圈是倾斜绕卷的。但是, Brenckle 在 1984 年 7 月检查模式标本后, 认为虽然切面不是非常好, 仍然可以推断出早期壳圈是平旋的 (Brenckle and Marchant, 1987; Brenckle, 1993, pl. 5, figs. 20-21)。正因为模式种与描述之间的不一致, 在后来的研究中就有了不同的意见。Miklukho-Maklay (1956) 将该种归入 *Planoarchaediscus*。其他学者有的认同模式种与描述之间的差异, 将该种放在 *Planoarchaediscus* 属 (如 Bogush and Yuferev, 1966; Brazhnikova in Brazhnikova *et al.*, 1967); 有的学者将其归入一个新的平旋的属 (如 Pirlet and Conil, 1977; Brenckle and Marchant, 1987)。本文采用后一种做法, 将该种归入 *Viseidiscus* 属。

**产地层位** 湖南邵阳马栏边剖面石磴子组 340.2-347 m (图3)。

**时代分布** 维宪期。俄罗斯、哈萨克斯坦、乌克兰、西欧、中国南部地区、土耳其、伊朗、埃及等地区。

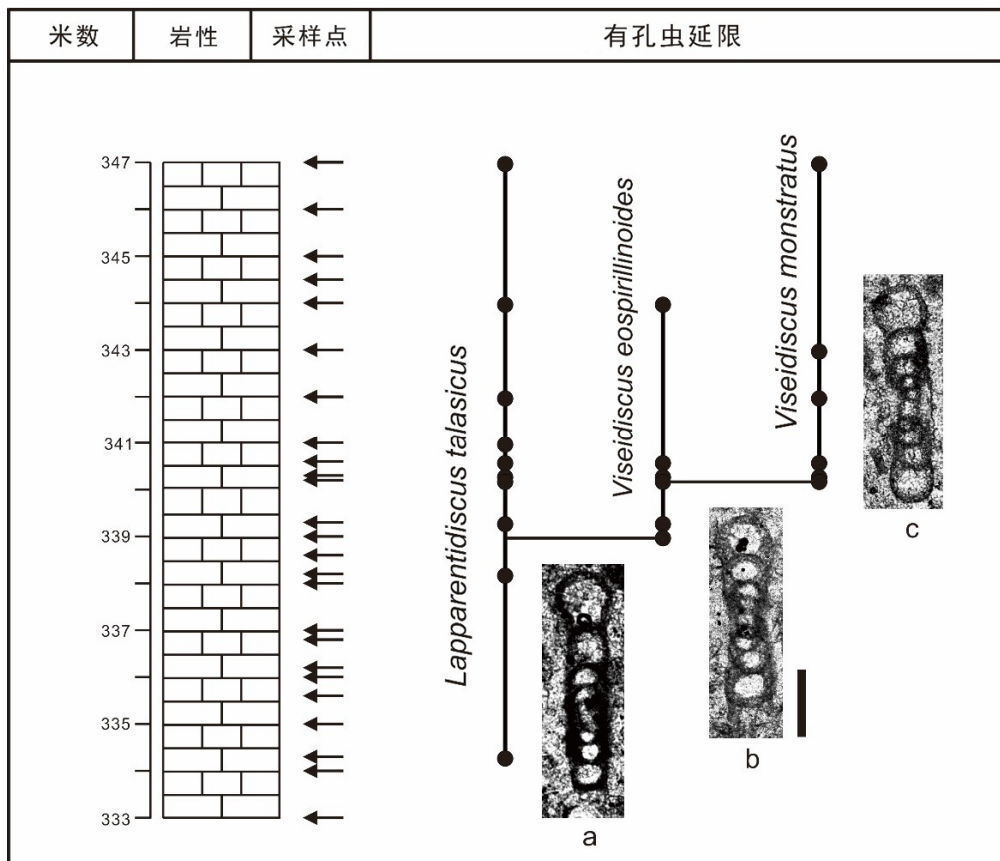


图3 马栏边剖面石磴子组中 *Lapparentidiscus* 与 *Viseidiscus* 的延限及演化图

Fig. 3 Range and evolution of *Lapparentidiscus* and *Viseidiscus* from the Shidengzi Formation at the Malanbian section  
 a. *Lapparentidiscus talasicus* (即图 2-5); b. *Viseidiscus eospirillinoides* (即图 2-8); c. *Viseidiscus monstratus* (即图 2-15)。比例尺为 100  $\mu\text{m}$ 。  
 a. *Lapparentidiscus talasicus* (see fig. 2-5); b. *Viseidiscus eospirillinoides* (see fig. 2-8); c. *Viseidiscus monstratus* (see fig. 2-15). Scale bar is 100  $\mu\text{m}$ .

#### 4 古盘虫类的早期演化序列

Conil等(1980)曾认为 *Brunsia* 和 *Pseudoammodiscus* 是古盘虫类的两个祖先, 而 Hance 等(2011)认为 *Lapparentidiscus* 是古盘虫类的唯一祖先。*Lapparentidiscus* 与 *Pseudoammodiscus* 的主要区别在于后者始终平旋、外旋, 壳圈匀速扩展; *Lapparentidiscus* 与 *Brunsia* 的主要区别在于后者早期壳圈绕旋。*Lapparentidiscus*、*Pseudoammodiscus* 和 *Brunsia* 在马栏边剖面石磴子组中都出现了, 但标本特征表明早期古盘虫类 *Viseidiscus* 的祖先应为 *Lapparentidiscus*。在石磴子组先后连续出现的一系列标本中, *Lapparentidiscus talasicus*、*Viseidiscus eospirillinoides*、*Viseidiscus monstratus* 这三个种在大小、形态、绕卷方式等方面非常相似, 都是早期平旋、内旋, 最后一两个壳圈外旋; 早期壳圈匀速扩

展, 最后一两个壳圈的扩展速度明显加快。这三个种最大的区别就在于它们的壳壁结构。*L. talasicus* 的壳壁是黑色微粒状; *V. eospirillinoides* 脐部的黑色微粒层中开始出现不连续的、斑驳的透明放射颗粒, 逐渐替代黑色微粒; *V. monstratus* 脐部的透明放射层已经比较发育、稳定和明显了, 有时还会延伸到最初的一两个壳圈的壳缘, 但是还未延伸到最后一两个壳圈的壳缘。这些标本展示了古盘虫类的早期演化序列(图3): *Lapparentidiscus talasicus* (Mikhno in Mikhno and Balakin, 1975)–*Viseidiscus eospirillinoides* (Brazhnikova, 1967)–*Viseidiscus monstratus* (Grozdilova and Lebedeva, 1954)。从而证实了两个观点: 1) *Lapparentidiscus* 是古盘虫类的祖先; 2) 古盘虫类的透明放射层是以替代而非填充的形式出现的, 透明放射层逐渐替代黑色微粒层, 并从脐部逐步扩展到壳缘。

## 5 结 论

华南马栏边剖面石磴子组先后连续出现的有孔虫标本证实了 *Lapparentidiscus* 是古盘虫类的祖先, 早期演化序列为 *Lapparentidiscus talasicus*–*Viseidiscus eospirillinoides*–*Viseidiscus monstratus*, 古盘虫类壳壁的黑色微粒层逐渐被透明放射层替代, 并从脐部扩展到壳缘。

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## 参考文献 (References)

- 湖南省地质矿产局区域地质调查队, 1987. 湖南晚泥盆世和早石炭世地层及古生物群. 北京: 地质出版社. 1–200.
- 林甲兴, 1984. 有孔虫目. 见: 宜昌地质矿产研究所(主编), 长江三峡地区生物地层学, 晚古生代分册(3). 北京: 地质出版社. 110–176.
- 林甲兴, 李家骧, 孙全英, 1990. 华南地区晚古生代有孔虫. 北京: 科学出版社. 1–297.
- 裘松余, 林甲兴, 1982. 湘中早石炭世晚期有孔虫群及其地层意义. 石油与天然气地质, 3: 177–183.
- 盛青怡, 2016. 华南密西西比亚纪有孔虫. 博士学位论文. 北京: 中国科学院大学. 1–285.
- 谭正修, 1990. 湖南石炭系底界及中间界线. 中国区域地质, 1: 51–59.
- 王克良, 1983. 湖南邵阳地区早石炭世的有孔虫. 中国科学院南京地质古生物研究所丛刊, 6: 209–224.
- Bogush O I, Bushmina E I, Domnikova O, 1970. О пограничных слоях девона и карбона Ельцовского синклиория в связи с изучением микрофауны – On the boundary layers of the Devonian and Carboniferous Yeltsovo synclinorium in connection with the study of microfauna. Труды института геологии и геофизики, Новосибирск: Наука – Proceedings of the Institute of Geology and Geophysics, Novosibirsk: Nauka, 71: 49–59 (in Russian).
- Bogush O I, Yuferev O V, 1966. Фораминиферы карбона и перми Верхояня – Foraminifera from the Carboniferous and Permian of Verkhoyan. Наука – Moscow: Izdat. “Nauka”. 1–209 (in Russian).
- Brazhnikova N E, Vakarchuk G I, Vdovenko M V, 1967. Микрофаунистические маркирующие горизонты каменноугольных и пермских отложений Днепровско-Донецкой впадины – Microfaunal reference horizons of Carboniferous and Permian deposits from the Dnieper-Donets basin. Киев: Наукова думка – Kiev, Naukova Dumka. 1–224 (in Russian).
- Brenckle P L, 1993. Type Archaetidiscacean foraminifers (Carboniferous) from the former Soviet Union and Great Britain. Cushman Foundation for Foraminifera Research Special Publication, 30: 1–59.
- Brenckle P L, Gaetani M, Angiolini L, Bahrammanesh M, 2009. Refinements in biostratigraphy, chronostratigraphy, and paleogeography of the Mississippian (Lower Carboniferous) Mobarak Formation, Alborz Mountains, Iran. GeoArabia, 14: 43–78. DOI: 10.2113/geoarabia140343
- Brenckle P L, Marchant T R, 1987. Calcareous microfossils, depositional environments and correlation of the Lower Carboniferous Um Bogma Formation at Gebel Nukhul, Sinai, Egypt. Journal of Foraminiferal Research, 17: 74–91. DOI: 10.2113/gsjfr.17.1.74
- Brenckle P L, Milkina N V, 2003. Foraminiferal timing of carbonate deposition on the late Devonian (Famennian)–Middle Pennsylvanian (Bashkirian) Tengiz Platform, Kazakhstan. Rivista Italiana Di Paleontologia E Stratigrafia, 109: 131–158.
- Brenckle P L, Ramsbottom W H C, Marchant T R, 1987. Taxonomy and classification of Carboniferous archaetidiscacean foraminifers. Courier Forschungsinstitut Senckenberg, 98: 11–24.
- Chernysheva N E, 1948. Об Archaediscus и близких к нему формах из нижнего карбона СССР – On Archaediscus and related forms from the Early Carboniferous of the USSR. In: Стратиграфия и фораминиферы нижнего карбона Русской платформы и Приуралья–Stratigraphy and Foraminifera of the Lower Carboniferous of the Russian Platform and Cis-Ural. Академия наук СССР Труды ИГН–Akademiya Nauk SSSR Trudy Instituta Geologicheskikh Nauk – Akademiya Nauk SSSR, Trudy Instituta Geologicheskikh Nauk, geologicheskaya seriya, 19(62): 150–158 (in Russian).
- Conil R, 1976. Contribution a l’etude des foraminifères du Dinantien de l’Irlande. Annales de la Société Géologique de Belgique, 99: 467–479.
- Conil R, 1980. Note sur quelques foraminifères du Strunien et du Dinantien d’Europe occidentale. Annales de la Société géologique de Belgique, 103: 43–53.
- Conil R, Longerstaey P J, Ramsbottom W H C, 1980. Matériaux pour l’étude micropaléontologique du Dinantien de Grande-Bretagne. Mémoires de l’Institut Géologique de l’Université de Louvain, 30: 1–187.
- Cushman J A, 1928. Foraminifera: their classification and economic use. Special Publication Cushman Laboratory for Foraminiferal Research, 1: 1–401. DOI: 10.5962/bhl.title.10175.
- Dain L G, Grozdilova L, 1953. Ископаемые фораминиферы СССР. Турнейеллиды и архедисциды – Fossil foraminifera of the USSR. Tournayellidae and Archaetidiscidae. Труды ВНИГРИ – Proceedings of the Oil Research Geological Institute (VNIGRI), 74: 1–115 (in Russian).
- Grozdilova L P, Lebedeva N S, 1954. Фораминиферы нижнего карбона и башкирского яруса среднего карбона Колво-Вишерского края–Lower Carboniferous, and Middle Carboniferous Bashkirian Foraminifera from Kolva-Vishera Region. Микрофауна СССР, Microfauna of the USSR, Proceedings of the Oil Research Geological Institute (ВНИГРИ-VNIGRI), 7: 4–236 (in Russian).
- Hance L, Hou H F, Vachard D, 2011. Upper Famennian to Visean foraminifers and some carbonate microproblematica from South



- China–Hunan, Guangxi and Guizhou. Beijing: Geological Publishing House. 1–359.
- Lin Jia-xing, 1984. Protozoa. *In*: Yichan Institute of Geology and Mineral Resources (ed.), *Biostratigraphy of the Yangtze Gorge area chiefly, Late Paleozoic era (3)*. Beijing: Geological Publishing House. 110–176 (in Chinese with English summary).
- Lin Jia-xing, Li Jia-xiang, Sun Quan-ying, 1990. Late Paleozoic foraminifers in South China. Beijing: Science Publishing House. 1–297 (in Chinese with English summary).
- Loeblich A R, Tappan H, 1984. Suprageneric Classification of the Foraminiferida (Protozoa). *Micropaleontology*, 30(1): 1–70. DOI: 10.2307/1485456
- Mamet B, 1975. *Viseidiscus*, un nouveau genre de Planoarchaediscinae (Archaeidiscidae, foraminifères). *Comptes Rendus Sommaires des Séances de la Société géologique de France*, 2: 48–49.
- Marfenkova M M, 1978. Фораминиферы и стратиграфия нижнего и среднего визе Южного Казахстана – Foraminifera and stratigraphy of Lower and Middle Visean from Southern Kazakhstan. *Тр. ИГиГ СО АН СССР – Proceedings of the Institute of Geology and Geophysics, Siberian Filiale of the Academy of sciences USSR*, 386: 78–99 (in Russian).
- Mikhno N M, Balakin G V, 1975. Фораминиферы и мшанки нижнего карбона Чаткальских гор – Foraminifera and bryozoans from the Early Carboniferous Chatkal horizon. Ташкент, ФАН Узбекской ССР – Ministerstvo Geologii Uzbekskoi SSR, Sredneasiatskii Institut Geologii i Mineralnogo Sviria, Izdatelstvo “Fan”, Uzbekskoy SSR: 1–54 (in Russian).
- Miklukho-Maklay A D, 1956. Биостратиграфическое расчленение верхнего палеозоя хр. Карачатыр (Южная Фергана) – Biostratigraphical divisions of the Upper Palaeozoic of the Kara-Chatir Ridge, S. Fergana. *Л ДАН СССР–CR Acad Sci URSS*, 108: 1152–1154 (in Russian).
- Miklukho-Maklay A D, 1957. Новые данные о систематике и филогении архедисцид – New data about the taxonomy and phylogeny of the Archaeidiscidae. *Вестник ЛГУ, Сер. геол. и геогр.* – *Bulletin of Leningrad State University, ser. geol. and geogr.*, 24: 34–46 (in Russian).
- Pirlot H, Conil R, 1977. L'évolution des Archaeidiscidae viséens. *Bulletin de la société belge de géologie de paléontologie et d'hydrologie*, 82: 241–299.
- Qiu Song-yu, Lin Jia-xing, 1982. The foraminiferal faunas of the later early Carboniferous in central Hunan and their stratigraphical significance. *Oil and Gas Geology*, 3: 177–183 (in Chinese with English abstract).
- Regional Geological Surveying Party, Bureau of Geology and Mineral Resources of Hunan Province, 1987. The late Devonian and early Carboniferous strata and Palaeobiocoenosis of Hunan. Beijing: Geological Publishing House. 1–200 (in Chinese with English summary).
- Shen Yang, Wang Xun-lian, 2015. Foraminiferal biostratigraphy of the Bei'an Formation (Visean–Serpukhovian) in the Pengchong area of Liuzhou, Guangxi, South China. *Alcheringa: An Australasian Journal of Palaeontology*, 39: 559–572. DOI: 10.1080/03115518.2015.1066579
- Sheng Qing-yi, 2016. Mississippian foraminifers from South China. Doctoral dissertation. Beijing: The University of Chinese Academy of Sciences. 1–285 (in Chinese with English summary).
- Tan Zheng-xiu, 1990. The lower boundary and intermediate boundary of the Carboniferous in Hunan. *Regional Geology of China*, 1: 51–59 (in Chinese with English abstract).
- Vachard D, 1980. Téthys et Gondwana au Paléozoïque supérieur; les données afghanes: biostratigraphie, micropaléontologie, paléogéographie. *Documents et Travaux IGAL, Institut Géologique Albert de Lapparent*, 2: 1–463.
- Vachard D, 1988. Towards a reasonable classification of the Archaeidiscidae (Foraminifera, Lower–Middle Carboniferous). *Revue de Paléobiologie*, 2: 103–123.
- Vdovenko M V, 1970. Найбільш давні (ранньовізейські) Ахедисцини великого Донбасу – The most ancient (Early Visean) Archaeidiscidae of the great Donbass. *Доповіді Академії наук української РСР – Reports of the Academy of Sciences of the Ukrainian SSR*, 12: 1061–1064 (in Russian).
- Wang Ke-liang, 1983. Early Carboniferous foraminifera from Shaoyang area of Hunan Province and their stratigraphic significance. *Bulletin of Nanjing Institute Geology and Palaeontology, Academia Sinica*, 6: 209–224 (in Chinese with English abstract).
- Wang Xiang-dong, Hu Ke-yi, Qie Wen-kun, Sheng Qing-yi, Chen Bo, Lin Wei, Yao Le, Wang Qiu-lai, Qi Yu-ping, Chen Ji-tao, Liao Zhuo-ting, Song Jun-jun, 2019. Carboniferous integrative stratigraphy and timescale of China. *Science China Earth Sciences*, 62: 135–153. DOI: 10.1007/s11430-017-9253-7
- Wang Xiang-dong, Qie Wen-kun, Sheng Qing-yi, Qi Yu-ping, Wang Yue, Liao Zhuo-ting, Shen Shu-zhong, Ueno Katsumi, 2013. Carboniferous and Lower Permian sedimentological cycles and biotic events of South China. *Geological Society, London, Special Publications*, 376: 33–46. DOI: 10.1144/sp376.11
- Yao Le, Qie Wen-kun, Luo Geng-ming, Liu Jiang-si, Algeo T J, Bai Xiao, Yang Bo, Wang Xiang-dong, 2015. The TICE event: Perturbation of carbon–nitrogen cycles during the mid-Tournaisian (Early Carboniferous) greenhouse–icehouse transition. *Chemical Geology*, 401: 1–14. DOI: 10.1016/j.chemgeo.2015.02.021.
- Zandkarimi K, Najafian B, Vachard D, Bahrammanesh M, Vaziri S H, 2016. Latest Tournaisian–late Visean foraminiferal biozonation (MFZ8–MFZ14) of the Valiabad area, northwestern Alborz (Iran): geological implications. *Geological Journal*, 51: 125–142. DOI: 10.1002/gj.2616

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